

Some new trends in applied mathematics: climate dynamics

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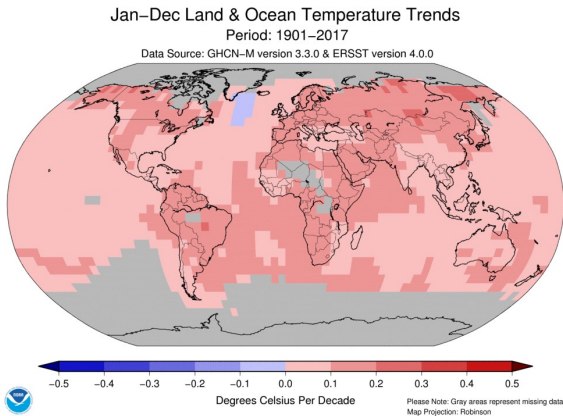
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Climate

Statistical long-term description of the quantities associated with weather (from months to millions of years)²

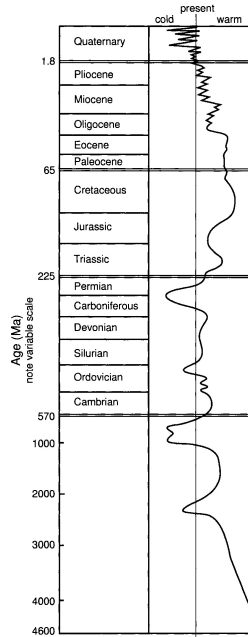
- Why?
 - To know the world.
 - To save the world.
- Who?
 - Climatologists.
 - Physicists.
 - **Mathematicians?!**
- Why?
 - Global measurements.
 - Climate proxy.
 - Mathematical models.



Palaeoclimatology

There are several ingenious ways to find what was the temperature millions of years ago!

- Climate proxy - indirect inference.
 - Ice cores (isotopes: ^{18}O and D (and many more)).
 - Sediments.
 - Tree rings and leaves.
 - Corals.
 - Foraminifera.
 - Pollen.
 - ...
- Figure: changes of global temperature.
Source: Saltzman, Dynamical Paleoclimatology (2002).



Palaeoclimatology

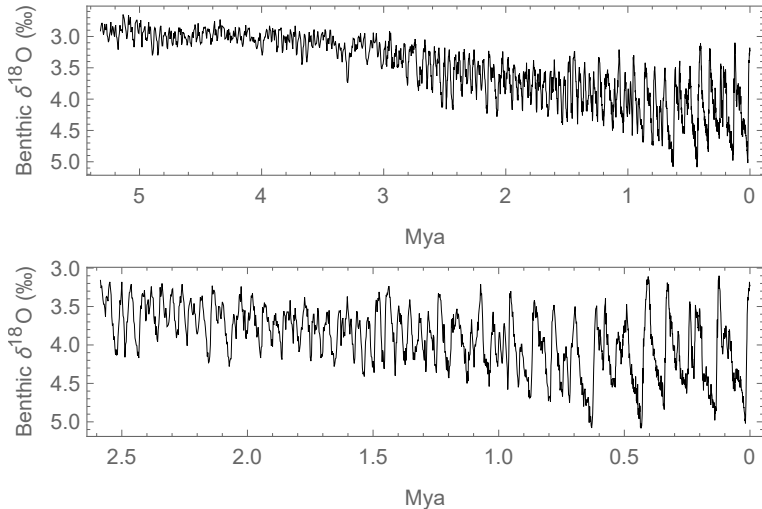


Figure: Climate proxy for the temperature. Source: Lisiecki and Raymo (2005).

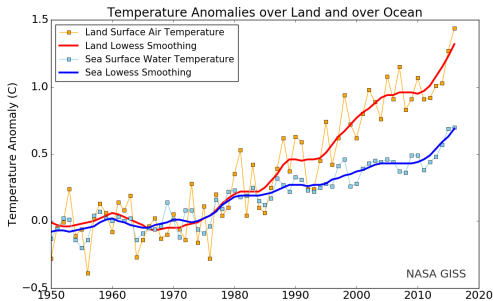
Ice ages

- Ice ages: huronian (early Proterozoic), cryogenian (late Proterozoic), andyjsko-saharyjskie (late Ordovician - Silurian), Karoo (Carbon) and quaternary (present).
- What is the reason for the climate oscillations-relaxations in the quaternary?
 - Astronomical forcing (Milankovitch theory?).
 - Changes in the atmosphere content?
 - Plate tectonics?
 - Internal nonlinear mechanisms?
- Snowball Earth
 - Hypothetical state of the climate in which almost the whole planet is under the ice (Neoproterozoic)
 - What is the mechanism responsible for initialization?
 - Positive feedback: colder → more ice → higher albedo → colder.
 - What is the mechanism responsible for termination? (volcanism?)

Human influence

*It is extremely likely that human influence has been the dominant cause of the observed warming since the mid-20th century.*³

- Humans emit enormous amounts of CO₂ - a greenhouse gas.
- The predicted increase in temperature in XXI century: 0.3 – 1.7°C (low emission); 2.6 – 4.8°C (highest emission).
- Consequences:
 - increase of global temperature,
 - raising sea level,
 - changes in the precipitation distribution,
 - extreme weather phenomena,
 - extinction of species.



Climate models

Complexity of hierarchy.

- Global Climate Models (GCMs)

- Geophysical Navier-Stokes on a sphere + equation of state + conservation of energy + conservation of salinity + other parametrisations

$$\rho \left(\frac{D\mathbf{u}}{Dt} + 2\boldsymbol{\Omega} \times \mathbf{u} \right) = -\nabla p + \nu \nabla^2 \mathbf{u} + \mathbf{f}, \quad \frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0,$$

$$p = p(\rho, T, s), \quad \frac{DT}{Dt} = q_T, \quad \frac{Ds}{Dt} = q_e.$$

- Earth System Models of Intermediate Complexity (EMICs)⁴

- GCM with more parametrisations/simplifications.

- Conceptual models

- Box models.
- **Dynamical systems.**

⁴Claussen, Martin, et al. „Earth system models of intermediate complexity: closing the gap in the spectrum of climate system models.” *Climate dynamics* 18.7 (2002): 579-586.

The simplest model

The simplest zero-dimensional model (Energy Balance Model)

- Let $T = T(t)$ be the globally averaged temperature. Then

$$c \frac{dT}{dt} = \underbrace{\frac{1}{4} (1 - \alpha(T)) Q}_{\text{incoming shortwave radiation}} - \underbrace{\gamma \sigma T^4}_{\text{outgoing longwave radiation}} =: f(T),$$

where c - specific heat, Q - solar constant, γ - greenhouse coefficient, σ - Stefan-Boltzmann constant and α - temperature dependent albedo.

- Albedo-temperature feedback: colder \rightarrow more ice \rightarrow larger albedo \rightarrow colder

$$\lim_{T \rightarrow 0^+} \alpha(T) = \alpha_-, \quad \lim_{T \rightarrow \infty} \alpha(T) = \alpha_+ \quad \text{oraz} \quad \alpha'(T) \leq 0,$$

where $\alpha_- > \alpha_+$.

Critical points

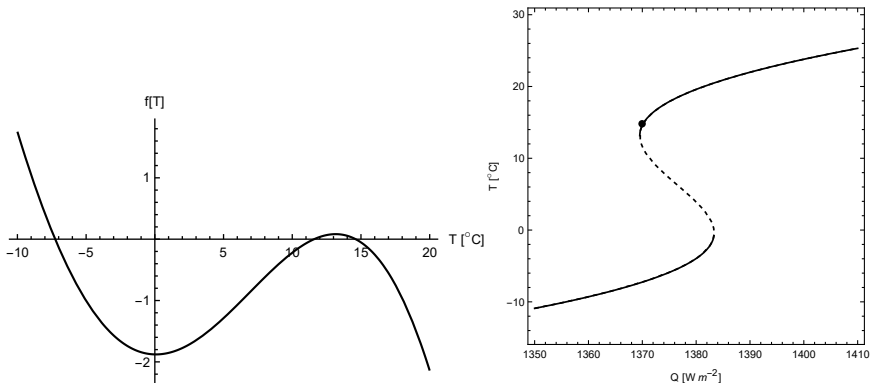


Figure: Exemplary simulation (data taken from A.C.Fowler, Mathematical Geoscience, 2011). On the left: zeros of f are the stationary points of the climate. On the right: bifurcation diagram with hysteresis.

Another dimension: Budyko-Sellers

- The previous model is based on the seminal papers of Budyko and Sellers^{5,6}.
- This time $T = T(y, t)$, where y is the sine of the latitude.
- Conservation of energy with the meridional transport

$$c \frac{\partial T}{\partial t} = \underbrace{Qs(y)(1 - \alpha(T))}_{\text{promieniowanie krótkofalowe}} - \underbrace{A + BT}_{\text{promieniowanie długofalowe}} - \underbrace{C(T - \bar{T})}_{\text{transport poziomy}},$$

where $s = s(y)$ is the meridional distribution of solar radiation, and \bar{T} is the mean temperature on Earth

$$\int_0^1 s(y) dy = 1, \quad \bar{T}(t) := \int_0^1 T(t, y) dy.$$

- North⁷ proposed diffusive transport.
- Modern and **mathematical** approach: McGehee and collaborators⁸.

⁵Budyko, M. I. (1969), *Tellus* 21(5), 611-619.

⁶Sellers, W. D. (1969), *Journal of Applied Meteorology* 8(3), 392-400.

⁷North, G. R. et al. (1981), *Reviews of Geophysics* 19(1), 91-121.

⁸McGehee, R. et al. (2012), *SIADS* 11(2), 684-707.

Further degrees of freedom

- Essential addition: **ice sheet extent**
 - At present: Greenland and Antarctica.
 - Last maximal glaciation: Laurentide, Patagonian and Vistulian.
- Model Källén, Crafoord, Ghil⁹ based on Weertman¹⁰

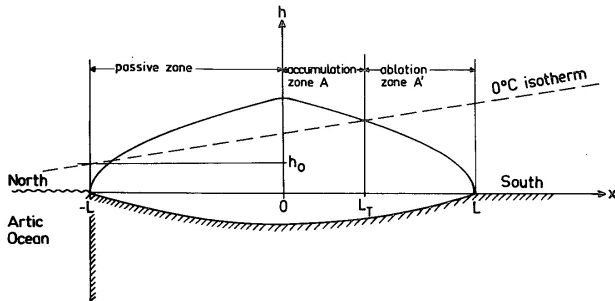


Figure: Ice sheet model according to KCG and Weertman.

⁹KCG (1979), Journal of the Atmospheric Sciences 36(12), 2292-2303.

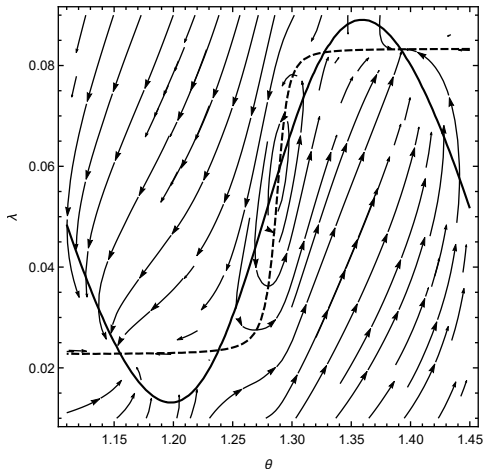
¹⁰Weertman, J. (1976), Nature 261, 17-20.

General KCG model

- KCG model implies internal climate oscillations (Hopf bifurcation).
- General KCG model **exhibits them also!** Is the climate an internal oscillator?

$$\frac{dV_I}{dT} = aL_T - m(L - L_T).$$

- Precipitation-temperature feedback: colder \rightarrow lower evaporation \rightarrow lower precipitation \rightarrow less ice \rightarrow lower albedo \rightarrow warmer.



What else?

- CO₂: the most important (but not the strongest) greenhouse gas.

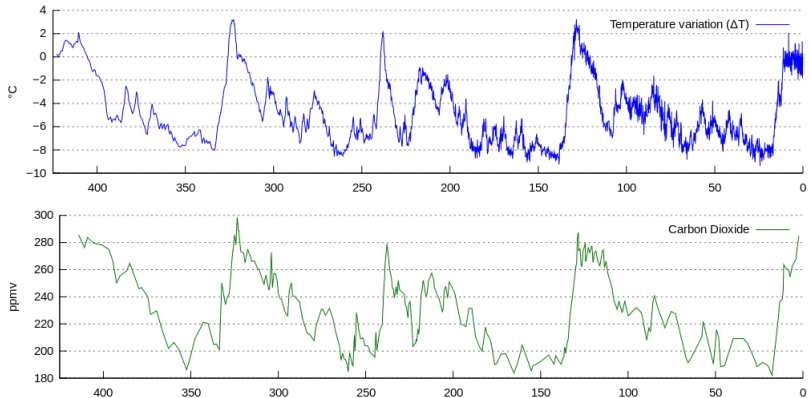


Figure: Historical CO₂ and the temperature at Vostok station.

Carbon cycle

- Fowler's „simple model”¹¹

$$\frac{dm_{CO_2}}{dt} = -A_L W(T, m_{CO_2}) + \nu,$$

where A_L - area of continents, W - weathering rate and ν - volcanism.

- Similarly to water, carbon has also its cycle:
 - Earth's mantle \rightarrow volcanoes (ν) \rightarrow atmosphere \rightarrow weathering (W) \rightarrow rivers \rightarrow oceans \rightarrow subduction.
- We need to couple that model with other equations¹²

$$W(T, m_{CO_2}) = W_0 m_{CO_2}^\mu \exp\left(\frac{T - T_0}{\Delta T}\right),$$

where $\mu \approx 0.3$.

¹¹Fowler, A. C., et al. (2013), International Journal on Geomathematics 4(2), 227-297.

¹²Walker, J. C. et al. (1981), Journal of Geophysical Research: Oceans, 86(C10),

The number of equations grows!

- What else?
 - The meaning of the oceans: CO₂.
 - We need to describe the chemistry of all oxidation reactions.
 - Mechanism for immediate entering into the interglacial period.
 - ...
- Model by Maasch and Saltzman¹³: ice mass, concentration of CO₂, temperature at the bottom of Atlantic Ocean.
 - Leading order of conservation laws.
 - Milankovitch + tectonic „trend”.
 - A change in the period of oscillations - bifurcation.
 - Mathematically: Bogdanov-Takens bifurcation¹⁴.
- Many more approaches...¹⁵
- **...and a lot of interesting mathematics!**

¹³M-S (1990), Journal of Geophysical Research: Atmospheres, 95(D2), 1955-1963.

¹⁴Engler, H. et al. (2017), arXiv preprint arXiv:1705.07387.

¹⁵Crucifix, M. (2012), Phil. Trans. R. Soc. A, 370(1662), 1140-1165.

Summary of the lecture

1. Introduction (now it ends!).
2. Budyko-Sellers energy balance model.
3. KCG precipitation-temperature feedback model.

THANK YOU!

Literature

■ This lecture is based on

1. Płociniczak, Ł. (2018). Hopf bifurcation in a conceptual climate model with ice-albedo and precipitation-temperature feedbacks. arXiv preprint arXiv:1801.09087.
2. Walsh, J., McGehee, R. (2013). Modeling climate dynamically. *The College Mathematics Journal*, 44(5), 350-363.
3. Fowler, A. (2011). *Mathematical geoscience* (Vol. 36). Springer Science & Business Media.

■ Models

1. Budyko, M. I. (1969). The effect of solar radiation variations on the climate of the earth. *tellus*, 21(5), 611-619.
2. Sellers, W. D. (1969). A global climatic model based on the energy balance of the earth-atmosphere system. *Journal of Applied Meteorology*, 8(3)
3. Källén, E., Crafoord, C., & Ghil, M. (1979). Free oscillations in a climate model with ice-sheet dynamics. *Journal of the Atmospheric Sciences*, 36(12), 2292-2303.

■ Textbooks on climate

1. Ruddiman, W. F. (2001). *Earth's Climate: past and future*. Macmillan.
2. Pierrehumbert, R. T. (2010). *Principles of planetary climate*. Cambridge